

Space-Time Regularizations for the ICON Model

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Outline

- 1 Introduction**
 - Motivation
- 2 3D Euler equations**
 - A 3D hydrostatic Euler model
 - Discretization
- 3 Numerical results**
 - cartesian geometry
 - spherical geometry
- 4 Advances in theory**
- 5 Outlook and open questions**



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Motivation for space- and time-regularization

- Mathematical models for atmospheric and oceanic flows exhibit the following problems
 - time step in explicit schemes is restricted by the fastest modes (e.g., gravity, acoustic waves)
Standard approach → **Semi-implicit time stepping schemes**
 - processes on unresolved (subgrid) scales interact in complex ways with those on resolved scales
Standard approach → **LES, RANS**
- Our approach → Introduce a **regularized set of analytical equations** which do not suffer from such problems



Motivation for space- and time-regularization

Our work considers two types of regularization

- **Time (τ) regularization** = Improve the computational efficiency by introducing a regularized (i.e. smoothed) pressure field
 - slows down the fastest waves
 - allows explicit time integration with a time step analogous to that of a semi-implicit method.
- **Space (α) regularization** = Parametrize the unresolved scales by introducing a regularized (i.e. smoothed) velocity field
 - conserves mass, potential vorticity, and potential enstrophy
 - preserves the validity of Kelvin's circulation theorem

Both regularized models

- can be interpreted as averaging out small scale fluctuations
- lead to analytic problems that can be proved to be well-posed
- do not enhance the viscosity of the flow



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3D Euler: The hydrostatic Euler equations

We consider a hydrostatic Euler system given by

$$\partial_t \mathbf{v} + (\mathbf{k} \cdot \nabla \times \mathbf{v} + f) \mathbf{v}^\perp + w \partial_z \mathbf{v} + c_p \theta \nabla \pi + \frac{1}{2} \nabla |\mathbf{v}|^2 = 0 \quad (1)$$

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \mathbf{v}) = 0 \quad (2)$$

$$\frac{\partial \mu}{\partial t} + \nabla \cdot (\mu \mathbf{v}) = 0 \quad (3)$$

augmented with a hydrostatic

$$c_p \theta \frac{\partial \pi}{\partial z} + g = 0 \quad (4)$$

and a thermodynamic relation

$$\frac{R}{p_s} \rho \theta = \pi^{\frac{1-\kappa}{\kappa}}. \quad (5)$$

Notation: $\mathbf{v} = (u, v)^T$ – horizontal velocity, w – vertical velocity, ρ – density, θ – potential temperature, c_p – specific heat of dry air, $\mathbf{k} = (0, 0, 1)^T$, g – free-fall acceleration, f – Coriolis parameter, p_s – surface pressure, R – gas constant, $\pi = \mu^{\frac{\kappa}{1-\kappa}}$ – Exner pressure, $\kappa \approx 0.4$.



3D Euler: τ -regularization

The τ -regularized system is defined by

$$\partial_t \mathbf{v} + (\mathbf{k} \cdot \nabla \times \mathbf{v} + f) \mathbf{v}^\perp + w \partial_z \mathbf{v} + c_p \theta \nabla \tilde{\pi} + \nabla \cdot \frac{1}{2} |\mathbf{v}|^2 = 0 \quad (6)$$

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \mathbf{v}) = 0 \quad (7)$$

$$\frac{\partial \mu}{\partial t} + \nabla \cdot (\mu \mathbf{v}) = 0, \quad (8)$$

where smoothed Exner pressure $\tilde{\pi}$ is computed from π by solving

$$\left[1 - \tau^2 \frac{c_s^2}{\rho \theta^2} \nabla \cdot (\rho \theta^2 \nabla) \right] \tilde{\pi} = \pi + \tau^2 \mathcal{R} \quad (9)$$

with $\tau \geq 0$, a smoothing parameter, and \mathcal{R} given by

$$\mathcal{R} = \frac{c_s^2}{c_p \rho \theta^2} \frac{\partial}{\partial z} \left(\frac{g \rho \theta}{1 + \tau^2 g \partial_z \log \theta} \right). \quad (10)$$

Notation: c_s – speed of sound



3D Euler: α -regularization

The α -regularized system is defined by

$$\partial_t \mathbf{v} + (\mathbf{k} \cdot \nabla \times \mathbf{v} + f) \tilde{\mathbf{v}}^\perp + w \partial_z \mathbf{v} + c_p \theta \nabla \pi \quad (11)$$

$$+ \nabla \left[\tilde{\mathbf{v}} \cdot \mathbf{v} - \frac{1}{2} \left(|\tilde{\mathbf{v}}|^2 + \alpha^2 |\nabla \tilde{\mathbf{v}}|^2 \right) \right] = 0 \quad (12)$$

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \tilde{\mathbf{v}}) = 0 \quad (13)$$

$$\frac{\partial \mu}{\partial t} + \nabla \cdot (\mu \tilde{\mathbf{v}}) = 0, \quad (14)$$

where $\tilde{\mathbf{v}}$ is computed from \mathbf{v} by solving

$$\left[1 - \alpha^2 \nabla \cdot \nabla \right] \tilde{\mathbf{v}} = \mathbf{v}, \quad (15)$$

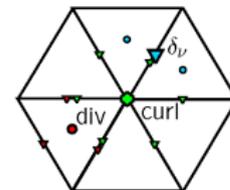
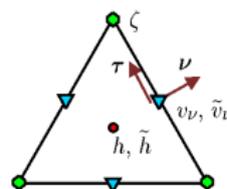
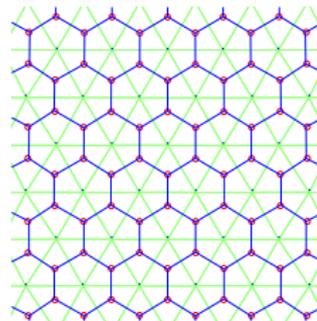
and $\alpha \geq 0$ is a smoothing parameter.



3D Euler: Horizontal discretization

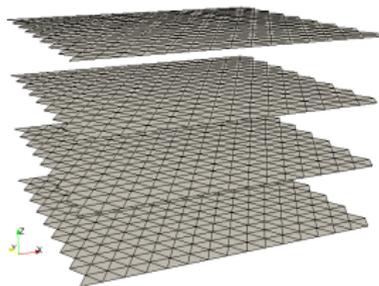
The horizontal discretization follows the ICON framework (Bonaventura, 2005) and is based on a triangular grid

- The normal velocities in the horizontal are located at edge midpoints
- Pressure and transported variables are located at the cell centers



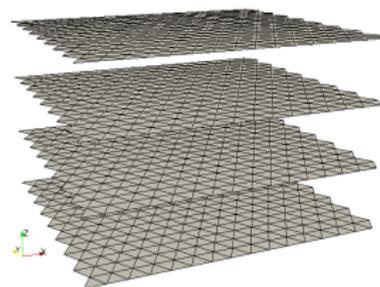
3D Euler: Vertical and time discretizations

- The vertical velocity w is located on layer interfaces.
- All variables except w are defined in the middle of a layer
- The vertical velocity is treated as a Lagrange multiplier that enforces the hydrostatic balance
- The regularization problem is solved in each horizontal layer; thus, we avoid solving a global 3D problem
- Several time stepping schemes are implemented
 - 3rd order total variation bounded (TVB) Runge-Kutta method (Cockburn et al, 1999)
 - 4th order strong stability preserving Runge-Kutta method (Spiteri, Ruuth, 2002)
 - Störmer-Verlet method



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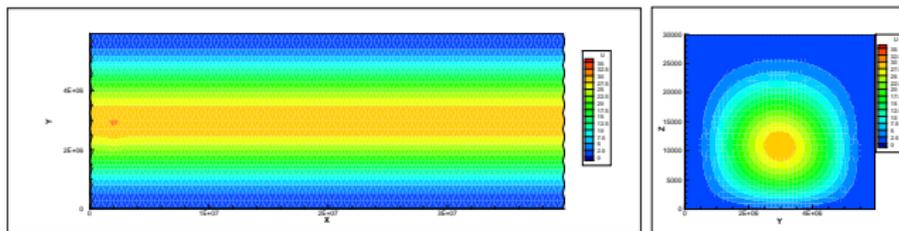
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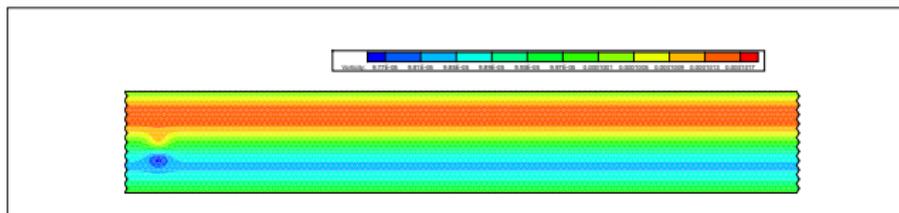
Baroclinic instability – cartesian case

The 3D model is run for the Jablonowski baroclinic instability test on an f-plane (Jablonowski, 2006)

- Domain: $[0, 40000] \times [0, 6000] \times [0, 30]$ (in km)
- Horizontal resolution of 312.5 km
- 30 horizontal layers with $\Delta z = 1000$ m
- The initial atmosphere is in hydrostatic and geostrophic balance except for a small wind perturbation
- Third order TVB Runge-Kutta method with Δt varying from 600s (unregularized) up to 2400s ($\alpha=1.25$)



Initial axial velocity; length-wise at $z = 11000$ m (left), crosswise at $x=20000$ km (right)



Initial vorticity at $z = 500$ m



Restoration of the geostrophic balance

The geostrophic balance equation

$$f\mathbf{v}^\perp + c_p \theta \nabla \pi = 0 \quad (16)$$

has for α -regularized systems the form

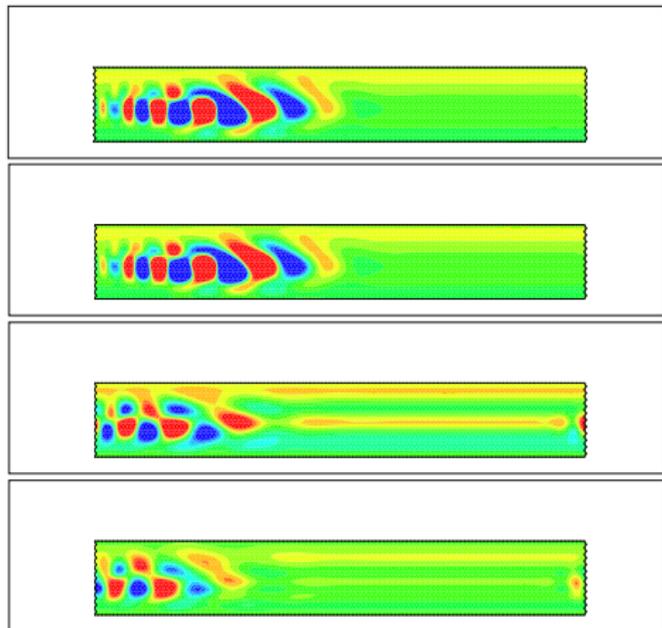
$$f\tilde{\mathbf{v}}^\perp + c_p \theta \nabla \pi = 0 \quad (17)$$

Possible solutions:

- Replace π by $\tilde{\pi}$ that satisfies the geostrophic balance
→ Requires knowledge of the analytical form of $\tilde{\mathbf{v}}$, delays advance of the baroclinic instability due to smoother initial pressure field
- Assign the balanced velocity to $\tilde{\mathbf{v}}$ and define \mathbf{v} using $\mathbf{v} = [1 - \alpha^2 \nabla \cdot \nabla] \tilde{\mathbf{v}}$
→ Can be computed automatically; produces a very rough initial velocity field for large α , difficult to stabilize.



Results(cartesian) – vorticity



Vorticity at $z = 500\text{m}$ at day 12. Top to bottom: no regularization, τ -regularization, α -regularization ($\alpha = 1$), α -regularization ($\alpha = 1.25$)

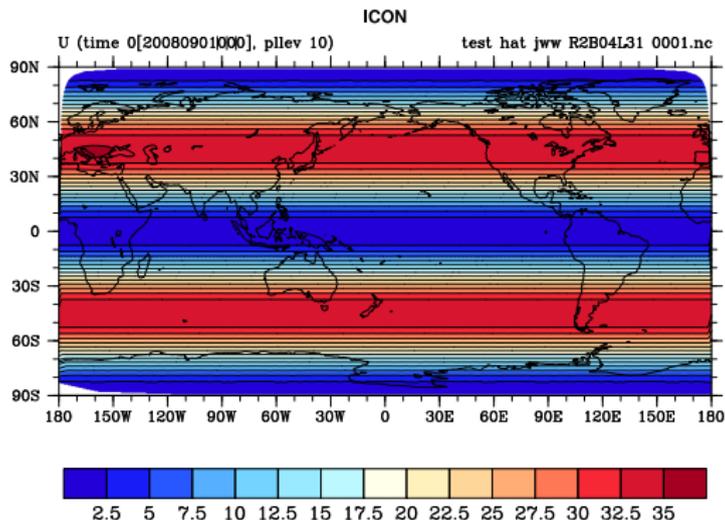
- The results of the τ -regularization model are very similar to those of the unregularized system
- α -regularization turns out to be very sensitive to the choice of alpha
- Initial conditions re-balanced by smoothing out the pressure field clearly slow down the advance of the instability



Baroclinic instability – spherical case

Jablonowski-Williamson baroclinic instability test on a sphere (Jablonowski, Williamson, 2007)

- Implementation within ICON hydrostatic core
- Initial condition re-balanced by applying the Helmholtz-operator to the horizontal velocity

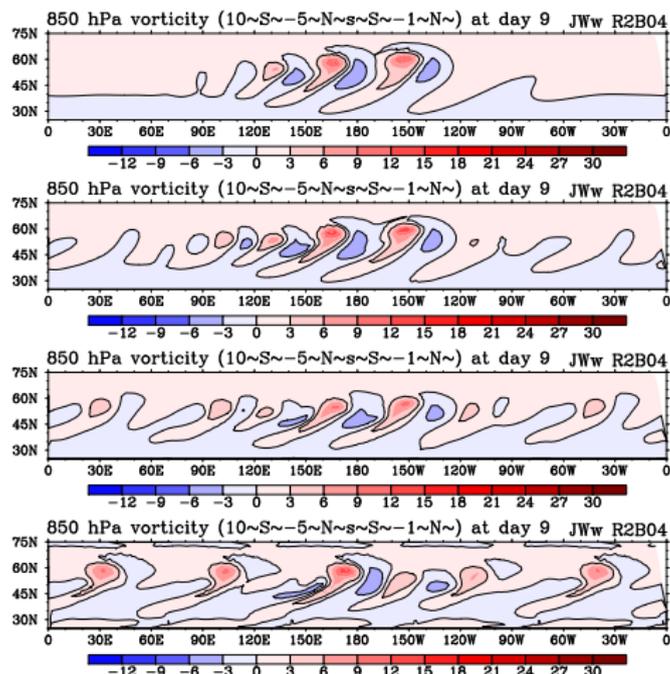


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Initial axial velocity (m/s) at 850hPa



Results(spherical) – vorticity



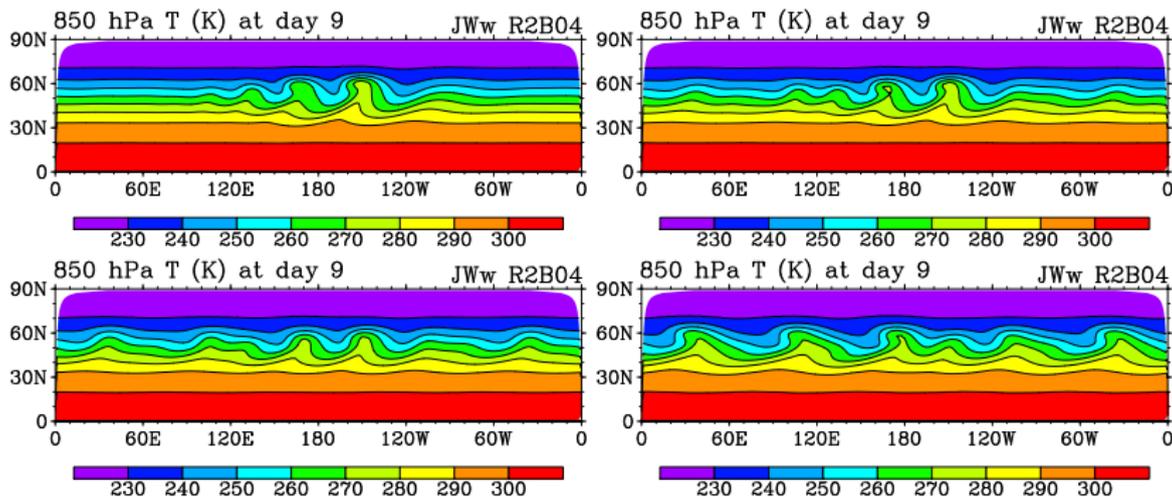
850hPa vorticity at day 9. Top to bottom: no regularization, α -regularization

($\alpha = 0.5$), α -regularization ($\alpha = 0.8$), α -regularization ($\alpha = 1$)

- Re-balancing of the initial conditions by applying the Helmholtz-operator to the velocity field increases the instability
- α -regularization allows for a much larger time steps than those of unregularized model



Results(spherical) – temperature



850hPa temperature at day 9. Top to bottom, left to right: no regularization, α -regularization ($\alpha = 0.5$), α -regularization ($\alpha = 0.8$), α -regularization ($\alpha = 1$)



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Theoretical results

Theoretical investigation of time and space regularization models conducted in the framework of this project resulted in

- a systematic derivation of a pressure regularization for the vertical slice Euler model in combination with a Störmer-Verlet method (Hundertmark and Reich, 2007)
- proof of the dependence of the longtime behavior of solutions to the Navier-Stokes- α , Leray- α , and Navier-Stokes- ω systems on a finite set of grid values or Fourier modes (Korn, 2011)
- an estimate of the number of determining nodes/modes in terms of flow parameters for each model (as above)



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Outlook and open questions

- *Boundary conditions*: The vast majority of theoretical and numerical work on the α -equations considers periodic domains. The extension to bounded domains poses questions such as:
 - How to choose the mathematically reasonable & physically sound boundary conditions?
 - How the smoothing parameter should be adapted in the vicinity of a boundary?
- *Non-uniform grids*: Grids with varying spatial resolution might need a scale-selective filtering in the transition region between different resolutions.
 - How the smoothing parameter α should be chosen to provide an adequate filter?
 - What are the dispersion properties of regularized models on such grids?

